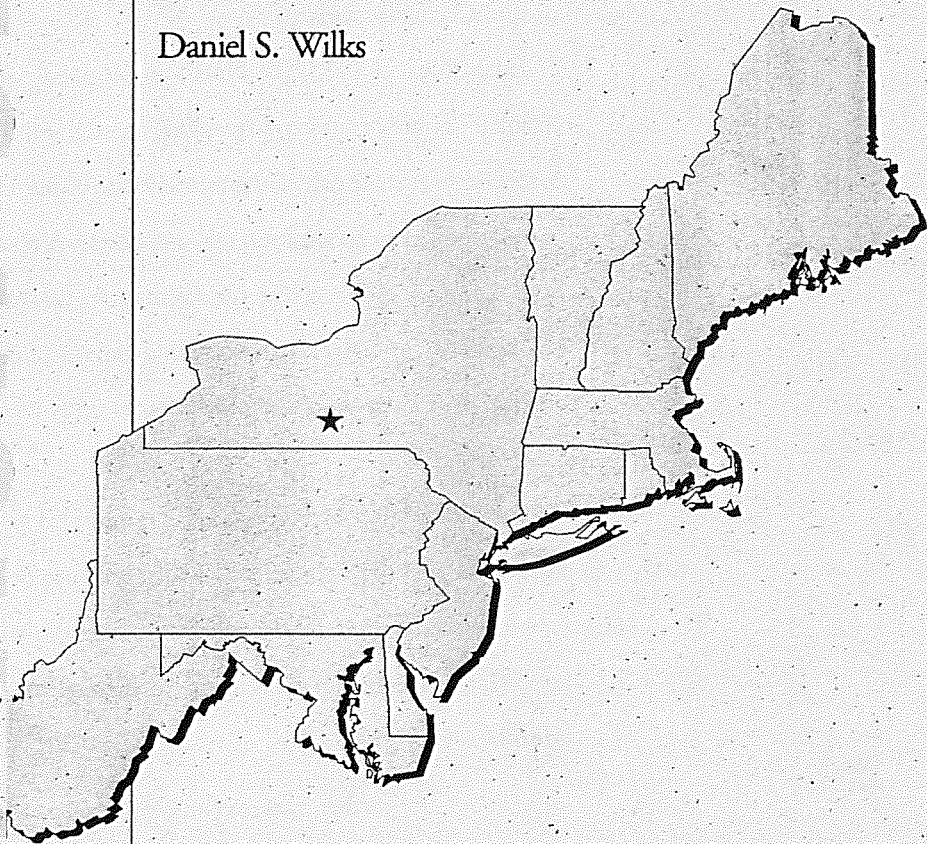


NORTHEAST REGIONAL CLIMATE CENTER

Spline-Interpolated Parameters for Adjusting Climatological Precipitation Distributions using the 30- and 90-Day Outlooks

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Ithaca, New York**

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INTRODUCTION

Recently a method has been described (Wilks and Eggleston 1992) which allows adjustment of gamma distributions representing monthly or seasonal precipitation totals in a manner consistent with the 30- and 90-day precipitation outlooks issued by the Climate Analysis Center (CAC) (Epstein 1988, Wagner 1989). The approach is analogous to that of Lehman (1987), which adjusts the parameters of Gaussian distributions representing average temperatures on the basis of the outlook probabilities for temperatures. The motivation in both cases is to allow users to evaluate probabilities for outcome thresholds other than the two (the 30th and 70th percentiles) to which the outlooks directly pertain.

STRUCTURE OF THE OUTLOOKS

The procedure is implemented through the conventional assumption that monthly or seasonal precipitation totals follow a gamma distribution, the probability density function for which is

$$f(x; \alpha, \beta) = \frac{(x/\beta)^{\alpha-1} \exp(-x/\beta)}{\beta \Gamma(\alpha)} ; x \geq 0; \alpha, \beta > 0 . \quad (1)$$

Here x is the random variable (total precipitation), α is the shape parameter, β is the scale parameter, and Γ denotes the gamma function. The function (1) is characterized by mean $\mu = \alpha\beta$ and variance $\sigma^2 = \alpha\beta^2$. It takes on a wide variety of shapes, ranging from strongly skewed and J-shaped for $\alpha \leq 1$ to approximately Gaussian for $\alpha \geq 50$. Using the gamma distribution is also convenient since monthly and seasonal precipitation are reported in terms of quantiles of climatological gamma distributions in the *Weekly Climate Bulletin* (Ropelewski et al. 1985).

The CAC outlooks divide the distribution of precipitation for a particular location and time period into three classes, separated by the 30th percentile (q_{30}) and the 70th percentile (q_{70}). These two points are given by

$$0.30 = \int_0^{q_{30}} f(x; \alpha_0, \beta_0) dx , \quad (2)$$

and

$$0.70 = \int_0^{q_{70}} f(x; \alpha_0, \beta_0) dx , \quad (3)$$

where the subscripts on the gamma distribution parameters denote that it is the unmodified climatological precipitation distribution being integrated. Precipitation outcomes smaller than q_{30} are denoted "dry," and outcomes larger than q_{70} are denoted "wet." Precipitation totals between these two quantiles are regarded as being "near normal."

The CAC outlooks specify a probability that the precipitation in the upcoming period will fall into either the "wet" or "dry" category, with the probability of "near-normal" always fixed at 0.4. The convention is adopted here and in Wilks and Eggleston (1992) that the forecast probability, z , pertains to the dry outcome. Thus the full forecast implied by a particular probability z is

$$\begin{aligned} \Pr(0 \leq x < q_{30}) &= z \\ \Pr(q_{30} \leq x < q_{70}) &= 0.4 \\ \Pr(q_{70} \leq x) &= 0.6 - z \end{aligned} \quad (4)$$

MODIFYING THE CLIMATOLOGICAL GAMMA DISTRIBUTION

The procedure is implemented by modifying the climatological distribution parameters α_0 and β_0 in a manner consistent with the outlook probability for precipitation below q_{30} . This is most conveniently done using the ratio of means, $\mu^*/\mu_0 = (\alpha^*\beta^*)/(\alpha_0\beta_0)$, and the ratio of scale parameters, β^*/β_0 , where starred symbols pertain to the modified distribution. These two ratios are related algebraically to the forecast z according to

$$\frac{\mu^*}{\mu_0} = 1 + b_M(z - 0.3) \quad (5)$$

and

$$\frac{\beta^*}{\beta_0} = 1 + b_1(z - 3) + b_2(z - 0.3)^2 \quad (6)$$

(Wilks and Eggleston 1992), which ensure that the modified and climatological distributions are identical when the climatological forecast $z=0.30$ is issued.

The three regression parameters b_M , b_1 , and b_2 in (5) and (6) vary as a function of the climatological shape parameter α_0 as shown in Figure 1. Use of this figure may be inconvenient if the procedure is to be used frequently, or is to be automated. Unfortunately, the relationships portrayed do not appear to have simple parametric representations. A practical alternative is to use cubic spline interpolation (e.g., Press et al. 1986) from the set of values presented in Table 1. This is accomplished implementing, for each of the three coefficients b in (5) and (6),

$$\hat{b} = A b_j + (1 - A) b_{j+1} + C b_j'' + D b_{j+1}'' \quad ; \quad (7)$$

where

$$A = \frac{\alpha_{j+1} - \alpha_0}{\Delta \alpha_j} \quad , \quad \alpha_j \leq \alpha_0 \leq \alpha_{j+1} \quad , \quad (8)$$

$$C = \frac{A^3 - A}{6} (\Delta \alpha_j)^2 \quad (9)$$

and

$$D = - \frac{A^3 - 3A^2 + 2A}{6} (\Delta \alpha_j)^2 \quad . \quad (10)$$

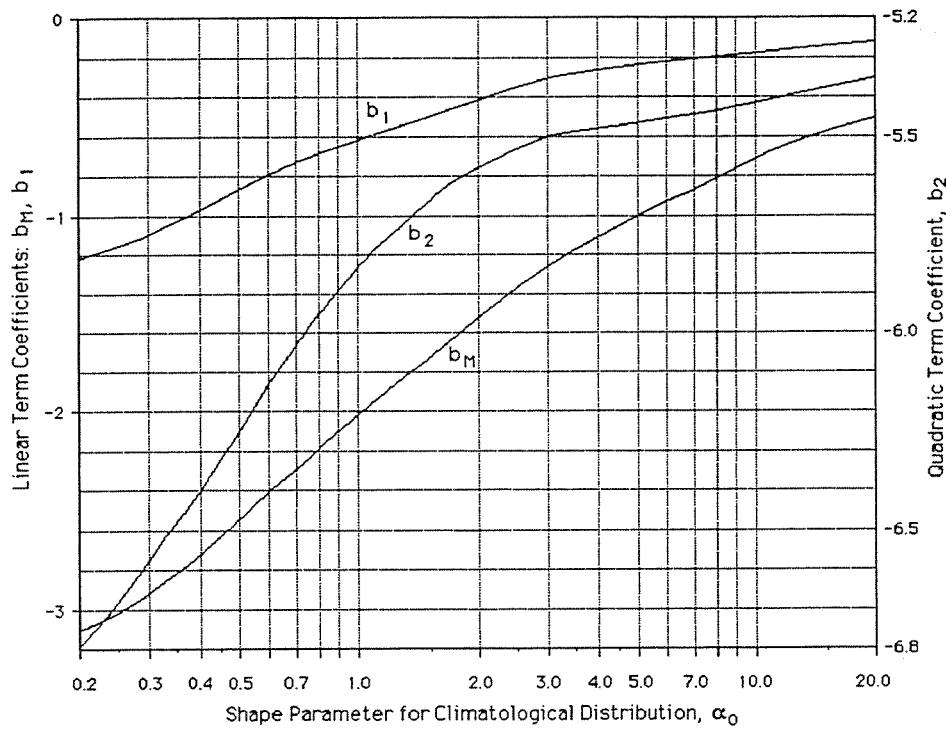


Figure 1. Dependence of the three parameters used in (5) and (6) on the shape parameter of the underlying climatological distribution, α_0 . Curves obtained using the spline interpolation algorithm (7) - (10) and the data in Table 1.

Table 1. Parameters for cubic spline interpolation of the parameters b_M , b_1 , and b_2 used in (5) and (6). Interpolation algorithm is given in (7) - (10).

j	α_j	$\Delta\alpha_j$	b_M	b''_M	b_1	b''_1	b_2	b''_2
1	0.1	0.1	-3.307	0.0000	-1.316	0.0000	-7.053	0.0000
2	0.2	0.3	-3.110	0.2619	-1.209	1.459	-6.795	-4.833
3	0.5	0.5	-2.553	-2.965	-0.870	-2.690	-6.252	-2.513
4	1.0	1.0	-2.035	-0.5169	-0.630	-0.0667	-5.833	-0.7241
5	2.0	3.0	-1.534	-0.1767	-0.425	-0.1047	-5.579	-0.0754
6	5.0	5.0	-1.010	-0.0091	-0.254	0.0054	-5.472	0.0056
7	10.0	10.0	-0.723	-0.0057	-0.172	-0.0033	-5.415	-0.0019
8	20.0	30.0	-0.515	-0.0004	-0.124	0.0002	-5.536	-0.0003
9	50.0	—	-0.327	0.0000	-0.074	0.0000	-5.391	0.0000

EXAMPLE

The procedure will be illustrated for the case of January total precipitation at Ithaca, New York. The available period of record is the 112 years 1879-1990, for which the fitted gamma distribution has $\alpha_0 = 4.41$ and $\beta_0 = 1.16$ cm. The parameters b_M , b_1 and b_2 are computed from Table 1 and (7)-(10). In this case, since $2.0 < \alpha_0 < 5.0$ (cf. equation (8)), $j=5$. Then,

$$A = (5.0 - 4.41)/3.0 = 0.1967, C = -0.2836, \text{ and } D = -0.4274.$$

Equation (7) then gives

$$b_M = A(-1.534) + (1-A)(-1.010) + C(-0.1767) + D(-0.0091) = -1.059.$$

Similarly, $b_1 = -0.260$ and $b_2 = -5.474$.

Consider the two (unusually confident) forecasts $z = 0.5$ and $z = 0.1$. In the former ("dry" forecast) case, using (6),

$$\beta^* = 1.16 \text{ cm} [1 + (-0.260)(0.5 - .3) + (-5.474)(0.5 - .3)^2] = 0.85 \text{ cm}.$$

Using (5),

$$\alpha^* = (\alpha_0) (\beta^*/\beta_0)^{-1} (\mu^*/\mu_0) = (4.41/0.729) [1 + (-1.059)(0.5 - .3)] = 4.77.$$

Similarly, for the "wet" forecast $z = 0.1$, $\beta^* = 0.97$ cm and $\alpha^* = 6.41$. These two adjusted gamma distributions are shown in Figure 4 of Wilks and Eggleson (1992).

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NOTES

NRCC RESEARCH SERIES

Knapp, W.W. and K.L. Eggleston, *Some Impacts of Recent Climate Variability on the Northeast*, NRCC Research Publication RR 91-1.

Wilks, D.S., *Gamma Distribution Probability Tables for Use in Climatology*, NRCC Research Publication RR 91-2.

Wilks, D.S., *Spline Interpolated Parameters for Adjusting Climatological Precipitation Distributions using the 30- and 90-Day Outlooks*, NRCC Research Publication RR 92-2.

NRCC DIGITAL DATA SETS

Eggleston, K.L. and D.S. Wilks, *Gridded Monthly Precipitation Distribution Parameters for the Continental United States*, NRCC Data Set DS 92-1.

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